Measurements and Modeling of Atmosphere-Snowpack Exchange of Ozone and Nitrogen Oxides at Summit, Greenland

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#### Why is the Arctic  $NO_x$  budget important?

 $NO<sub>2</sub> + OH + M \rightarrow HNO<sub>3</sub> + M$ 

Is this the termination of  $NO<sub>x</sub>$ ?



Maximum extent of snow cover in Northern Hemisphere Retrieved on 10/6/13 http://nsidc.org/cryosphere/sotc/snow\_extent.html



Honrath, J. Geophys. Res, 2000

## Summit, Greenland

- Two year campaign
- Elevation 3km above sea level
- Over 300 km to ocean
	- Minimal halogen chemistry
- Pollution controlled camp
	- Electrical vehicles





#### Met and Snow Towers



#### Seasonal Trends in the Measurements



#### Chemical measurements



#### Physical measurements

![](_page_6_Figure_2.jpeg)

#### Temperature Wind Direction and Speed

![](_page_6_Figure_4.jpeg)

## 1-D Process-scale snowpack model

- Goals
	- Reproduction of observed chemical trends in snowpack
	- $-$  Creation of a simplified model for estimation of NO<sub>x</sub> fluxes from snowpack.
- Overview of Components
	- Physical representation Snow density, porosity, snowflake radius, aqueous phase (QLL) on surface of snowflakes
	- Chemistry Gas and Aqueous phase
	- Physical Transport Diffusion, wind pumping, and mass transfer between phases

## Physical Representation of Snow

$$
SSA = -308.2 * LGG(\rho_{snow}) - 205.96
$$

$$
r = \frac{3}{SSA * \rho_{ice}}
$$

$$
por = 1 - \frac{\rho_{snow}}{\rho_{ice}}
$$

- $\rho_{\text{snow}}$  density of snow (0.3-0.6 g/cm<sup>3)</sup>
- SSA specific surface area  $\text{cm}^2/\text{gl}$
- $r$  radius of snowflake
- por porosity of snowpack

Domine, Atmos. Chem. Phys ,2008

![](_page_8_Figure_7.jpeg)

# Chemistry

![](_page_9_Picture_22.jpeg)

- Photolysis rates calculated with Fast-JX
- Chemistry calculated with Kinetic Pre-Processor (KPP)

## Physical Transport Part1 Wind pumping and Diffusion

$$
\frac{\partial c_g}{\partial t} = -\nabla \cdot \left( U_{firm} c_g \right) + \nabla \cdot \left( D_g \nabla c_g \right)
$$
\n
$$
U_{firm} = \frac{6k\rho_{air}}{\pi \mu \lambda_{surf}} \frac{h}{\lambda_{surf}} \frac{\sqrt{\alpha^2 + 1}}{\alpha} u_{10}^2 \left( C_1 \exp\left(\frac{z}{\delta}\right) \right)
$$
\n
$$
- C_2 \exp\left(\frac{z}{\delta}\right)
$$
\n
$$
\delta = \frac{1}{2} \frac{\alpha}{\sqrt{\alpha^2 + 1}} \frac{\lambda_{surf}}{\pi}
$$
\n
$$
C_1 = \frac{\exp\left(\frac{H_s}{\delta}\right)}{\exp\left(\frac{H_s}{\delta}\right) + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_2 = \frac{\alpha}{\pi \mu \sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_3 = \frac{\alpha}{\pi \sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_4 = \frac{\alpha}{\pi \sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_5 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_6 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_7 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_8 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_9 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_1 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_3 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_4 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)} \qquad C_5 = \frac{\alpha}{\sqrt{\frac{H_s}{\delta}} + \exp\left(-\frac{H_s}{\delta}\right)}
$$

- $U_{\text{firm}}$  vertical wind speed in snow
- $u_{10}$  wind speed 10 meters high
- $k$  permeability of snow
- $\mu$  dynamic viscosity of air
- $\lambda_{\text{surf}}$  relief wavelength
- h relief amplitude
- $\rho_{\text{air}}$  density of air
- $\alpha$  horizontal aspect ratio
- $D_g$  Gas diffusion coefficient
- $c_g$  Gas phase concentration
- $H_s$  Depth of ventilated snow

Toyota and McConnell via Liao, "Atmos. Chem. Phys, 2008

![](_page_10_Figure_14.jpeg)

## Physical Transport Part2 Mass Transfer

$$
\frac{dc_g}{dt} = -L\frac{dc_a}{dt} = -Lk_{mt}\left(c_g - \frac{c_a}{K_H}\right)
$$

$$
k_{mt} = \left(\frac{r^2}{3\ D_g} + \frac{4r}{3\nu\alpha}\right)^{-1}
$$

 $c_g$  – Gas phase concentration

 $c_a$  – Aqueous phase concentration

L – Volumetric ratio aqueous/gas phases

 $k<sub>mt</sub>$  – Mass transfer coefficient

 $K_H$  – Henry's Law constant [dimensionless]

r – radius of snowflake

- $D_g$  Gas phase diffusion coefficient
- $\alpha$  accommodation coefficient

v – molecular velocity

Sander , Surveys in Geophysics ,1999

Comparison of Model to Measurements

- Date of comparison is  $4/11/09 4/25/09$ – Ozone intrusion event
- Model nudged with surface measurements of  $NO<sub>x</sub>$ ,  $O<sub>3</sub>$ , temperature, and wind speed

## Model vs. Measurements

 $\mathsf{U}_3$ 

![](_page_13_Figure_1.jpeg)

# Model vs. Measurements  $NO<sub>2</sub>$

![](_page_14_Figure_1.jpeg)

#### Model vs. Measurements NO

![](_page_15_Figure_1.jpeg)

# Conclusion

- Model results
	- Over predication of O3
	- Over/under prediction of NO
	- $-$  Trend of NO<sub>2</sub> present, but not identical
- Major uncertainties in model
	- Does the QLL really behave as an aqueous phase?
	- What is the sensitivity of wind pumping to micro-topography parameters?
	- Is it acceptable to model snowflakes as spherical?
- Looking forward
	- Identify major chemical pathways
		- New laboratory experiments reveal relative humidity could significantly affect release of HONO from QLL [Finlayson-Pitts, UC Irvine, 2013]
	- $-$  Calculate NO<sub>x</sub> fluxes from snowpack
	- Combine major components into simplified model for possible integration into global models.

## Acknowledgments

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![](_page_17_Picture_5.jpeg)

#### References

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Questions?

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